

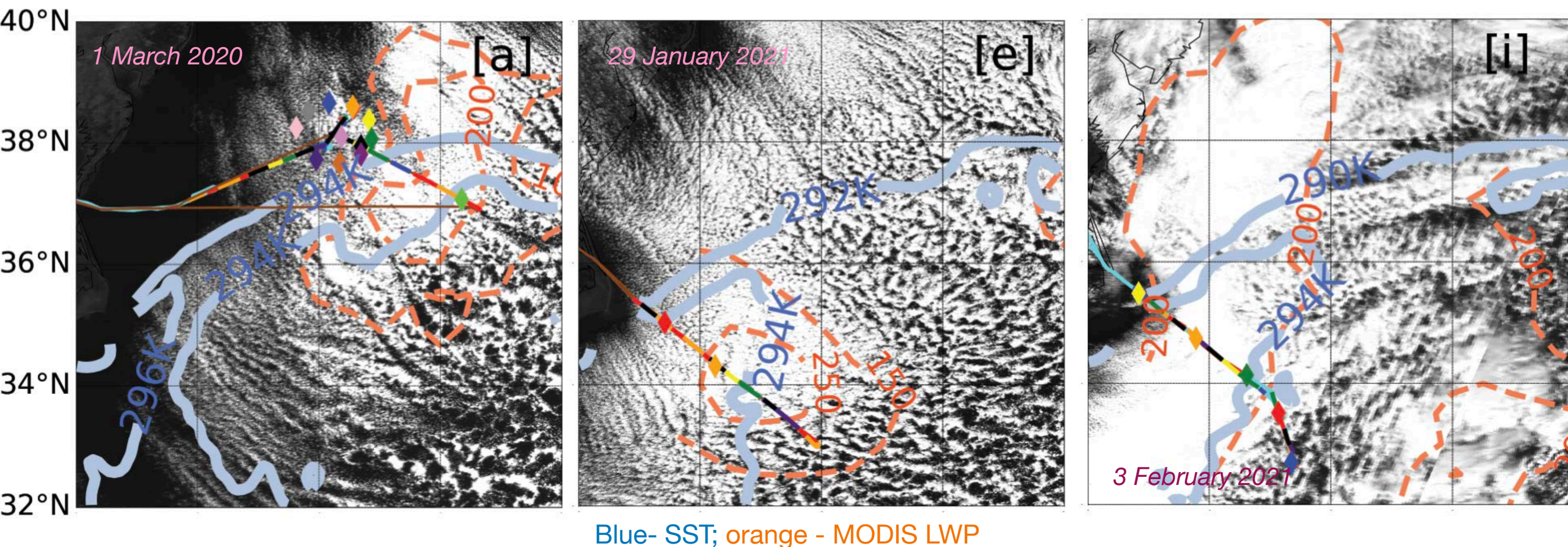
University of Miami ACTIVATE effort  
Seethala Chellappan, Paquita Zuidema

1. Examine the winter cold-air outbreak data from the point of view of generalizing their attributes & extending to satellite, reanalysis datasets. Focusing on 5 cases we believe encompass the phase space in  $N_d$ , LWP,  $T_{ct}$
2. This includes assessing remotely-sensed cloud properties ( $N_d$ , LWP,  $T_{ct}$ )
3. Examine cloud microphysics as a function of ASTER-derived cloud spatial scales

## What we did

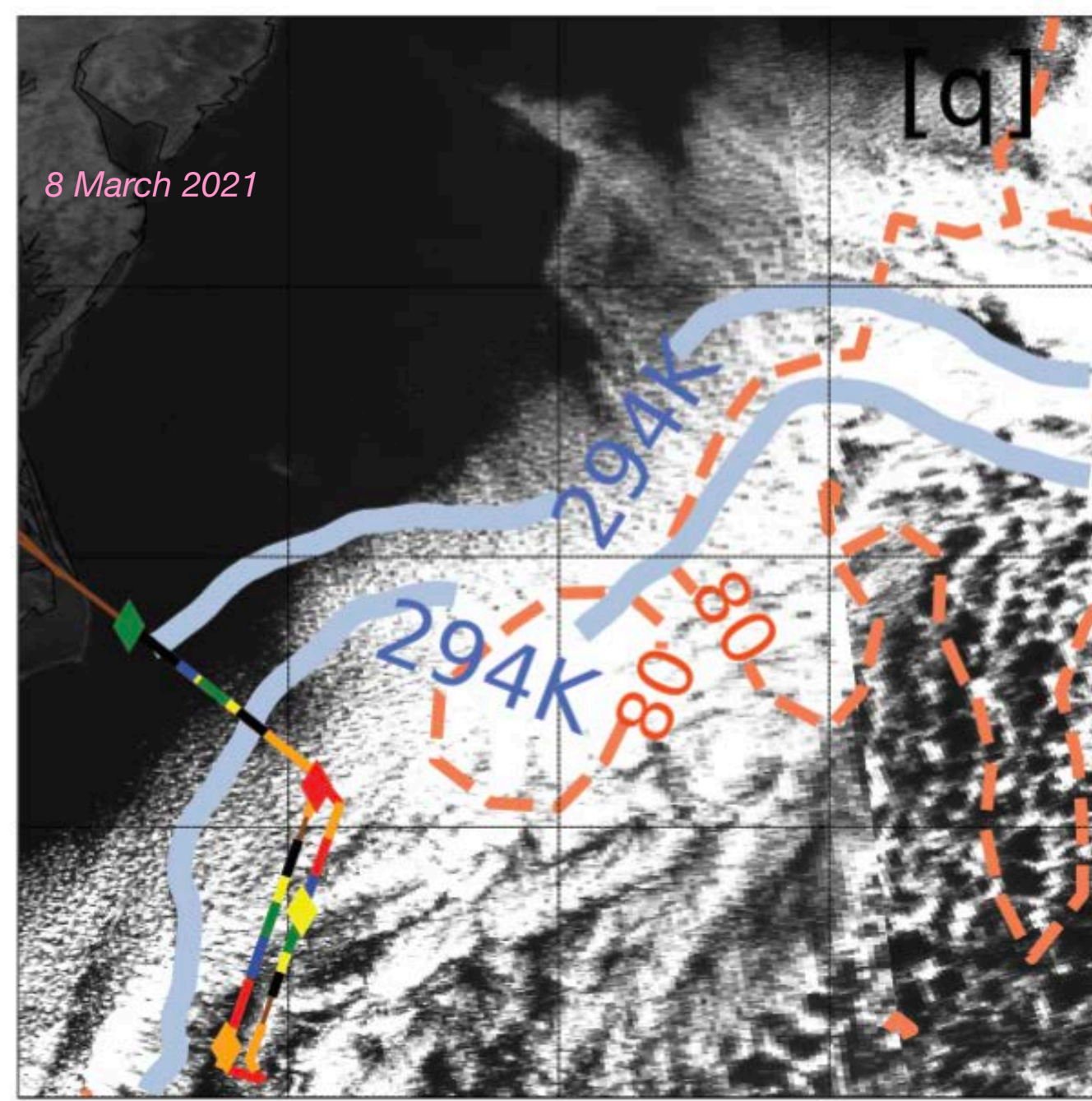
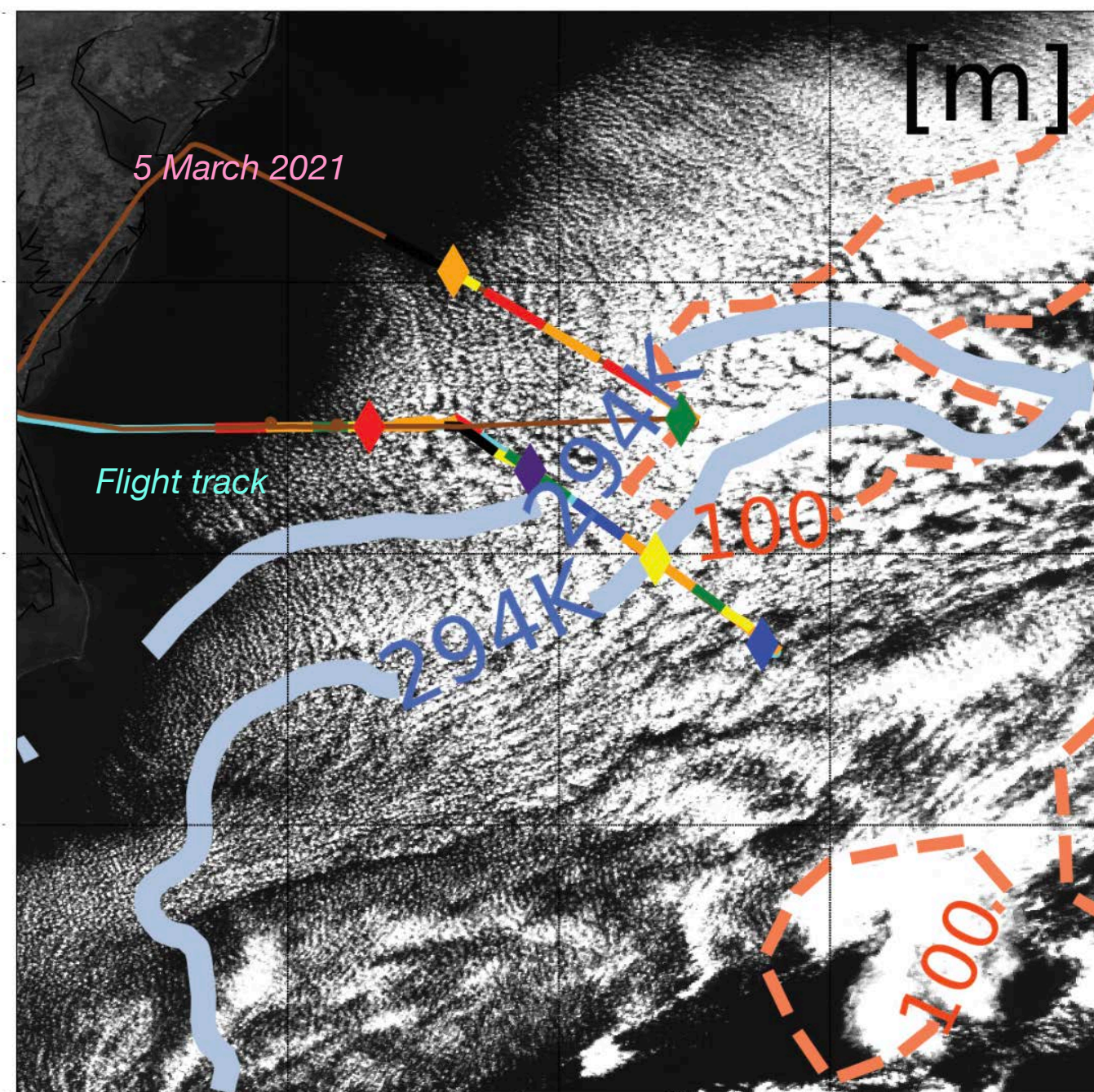
- Selected 5 cold-air-outbreak cases spanning a representative range of environmental/aerosol conditions
  - flight-maxima LWPs: 80-250 g m<sup>-2</sup> (MODIS)
  - cloud top temperatures flight-minima: -5 to -15 °C
  - in-situ* flight-maxima N<sub>d</sub> \*: 500-1540 cm<sup>-3</sup>

=> examine the *in-situ* ice microphysics for dependencies  
***a framework***



\*continental aerosol outflow; *Painemal et al 2021 JGR*; *Tornow et al 2021 GRL*.  
*in-situ* near-cloud-top effective radii only reach 9 micron at best.





Blue- SST; orange - MODIS LWP



What we have learned #1:

## **Wintertime cold-air outbreak clouds over the western Atlantic are not ice-deprived**

4 out of the 5 cases already contained ice as soon as clouds developed

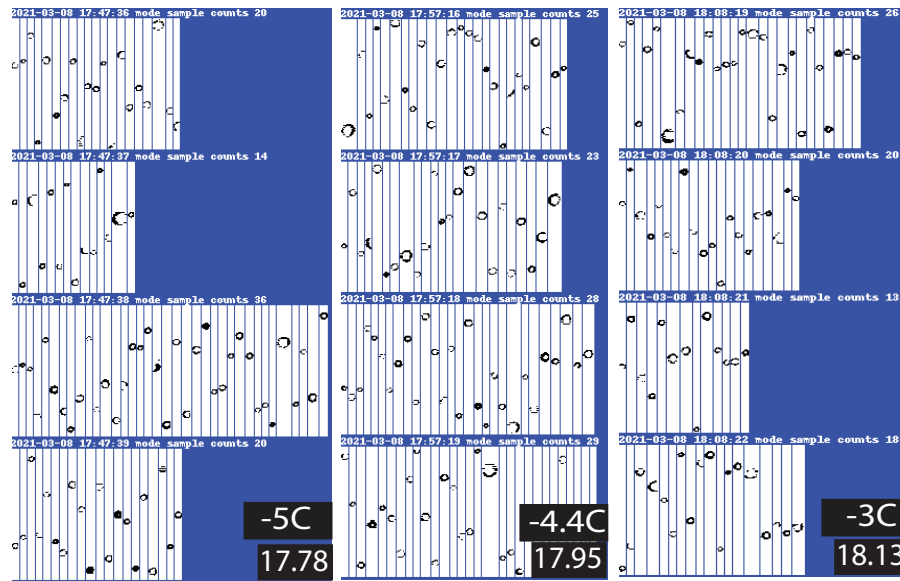
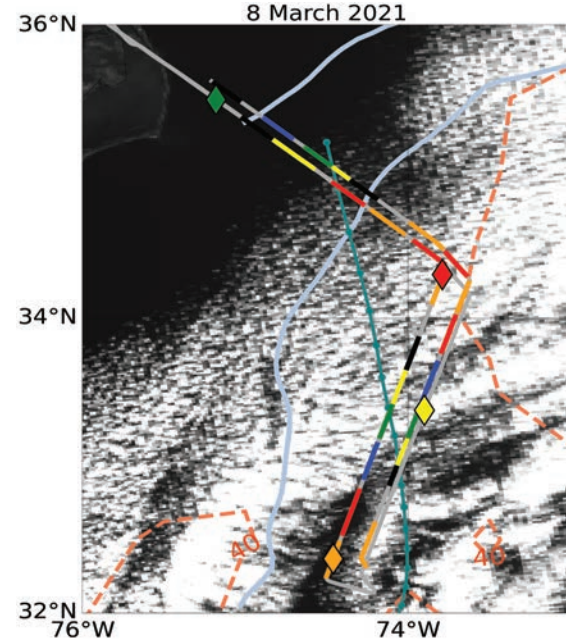
This despite cloud top temperatures  $> -15^{\circ}\text{C}$ , and small droplet sizes

Original premise that clouds start all-liquid then transition to mixed-phase  
thrown out the window

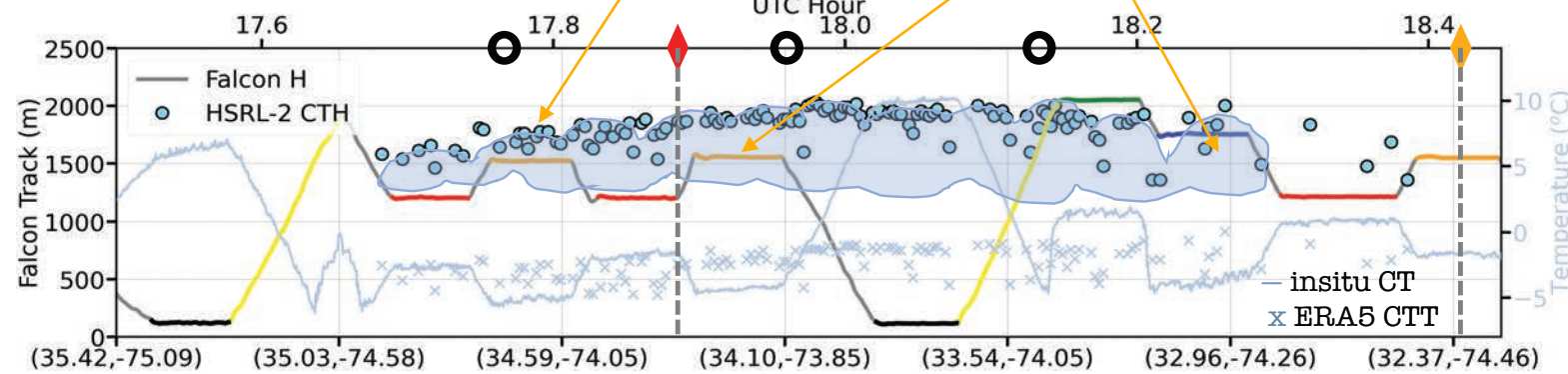
# 8 March 2021 (afternoon)

[a] MODIS visible radiance

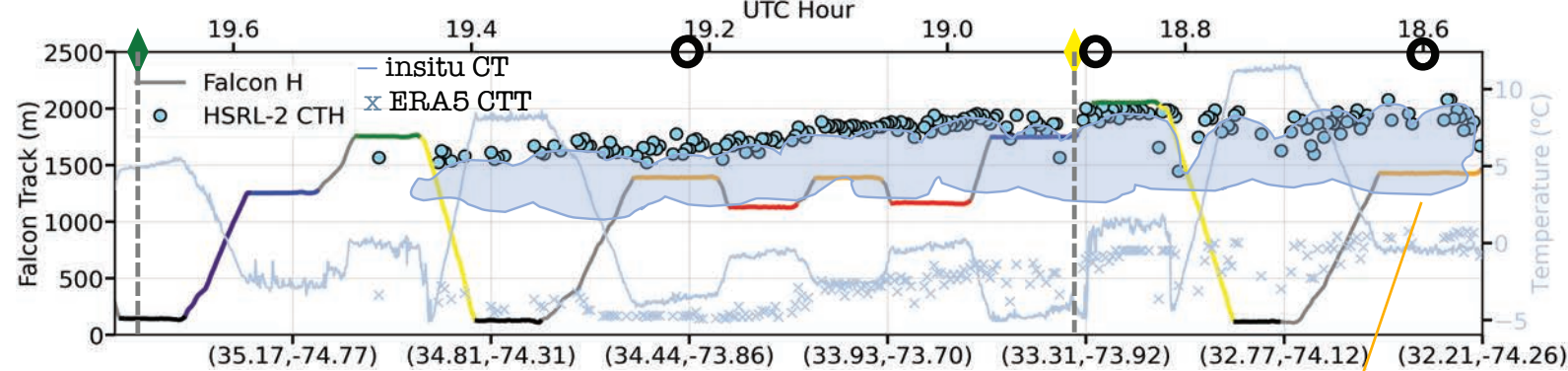
[b] 2D-Stereo images along Rf51 outbound Flight leg in [c]



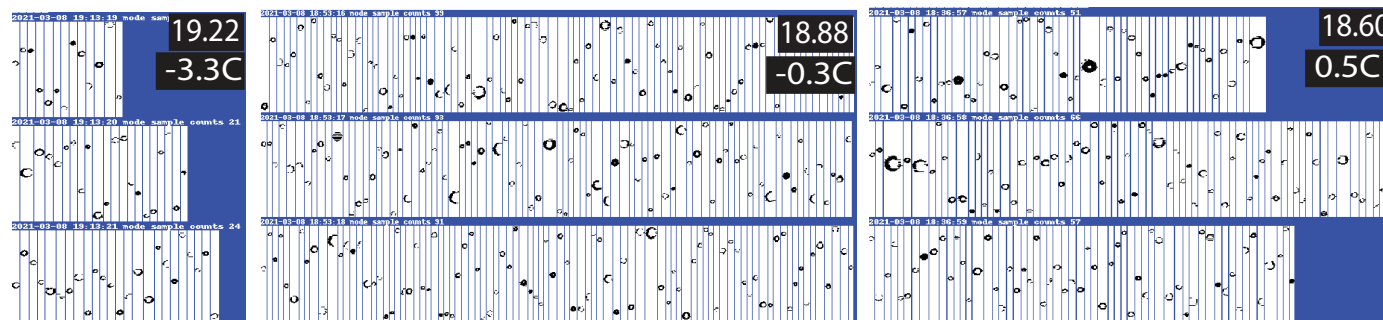
[c] Rf51 outbound Flight leg



[d] Rf51 inbound Flight leg



[e] 2D-Stereo images along Rf51 inbound Flight leg shown in [d]



*the one flight with  
no evidence of ice*

$$\overline{N}_d = 1000 - 1100 \text{ cm}^{-3}$$

$$T_{ct} \sim -5^\circ\text{C}$$

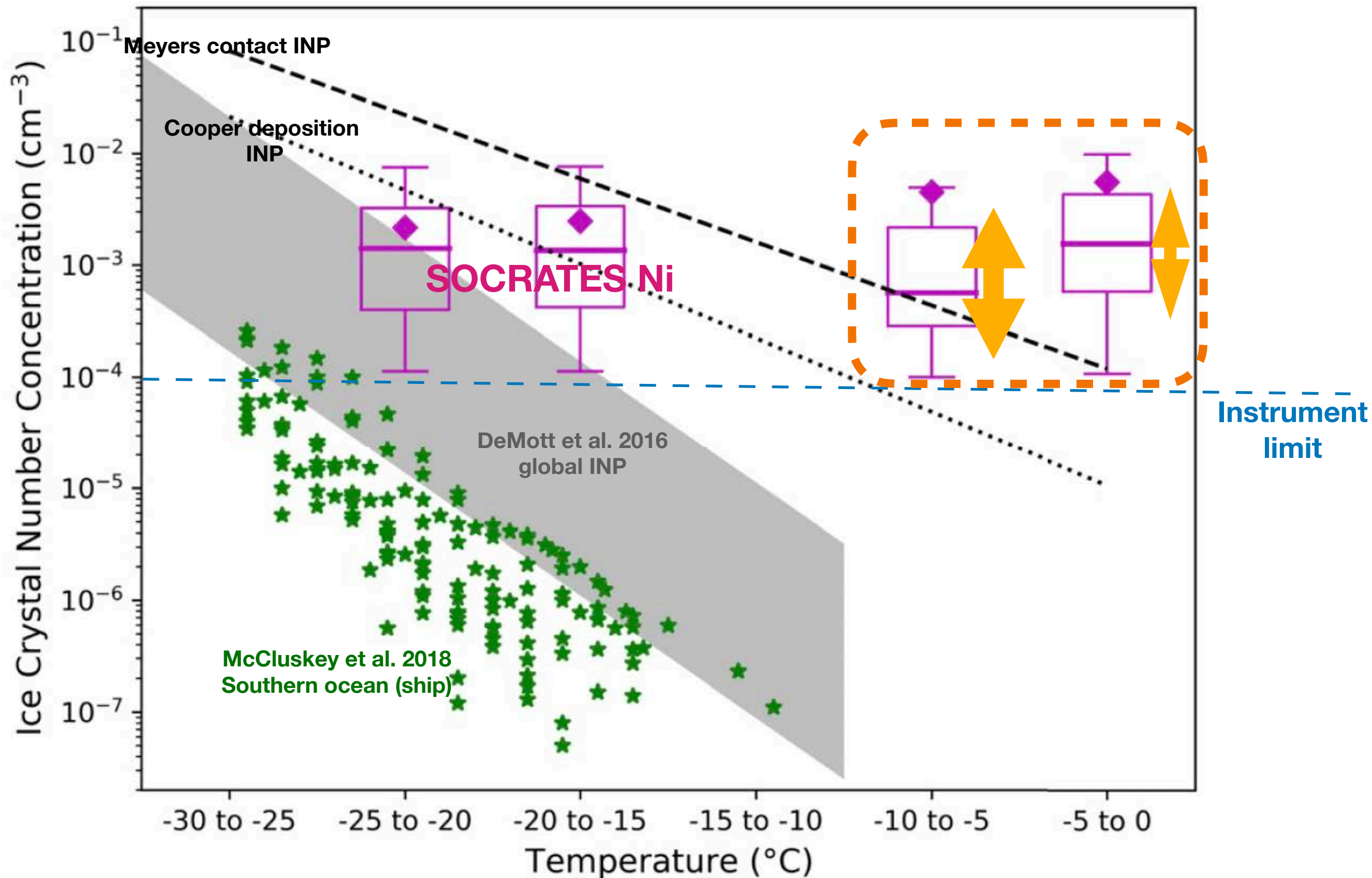
$$\text{RSP LWPs} < \sim 80 \text{ g m}^{-2}$$

Surface wind speeds low ( $\sim 8 \text{ m/s}$ )

No near-surface rain

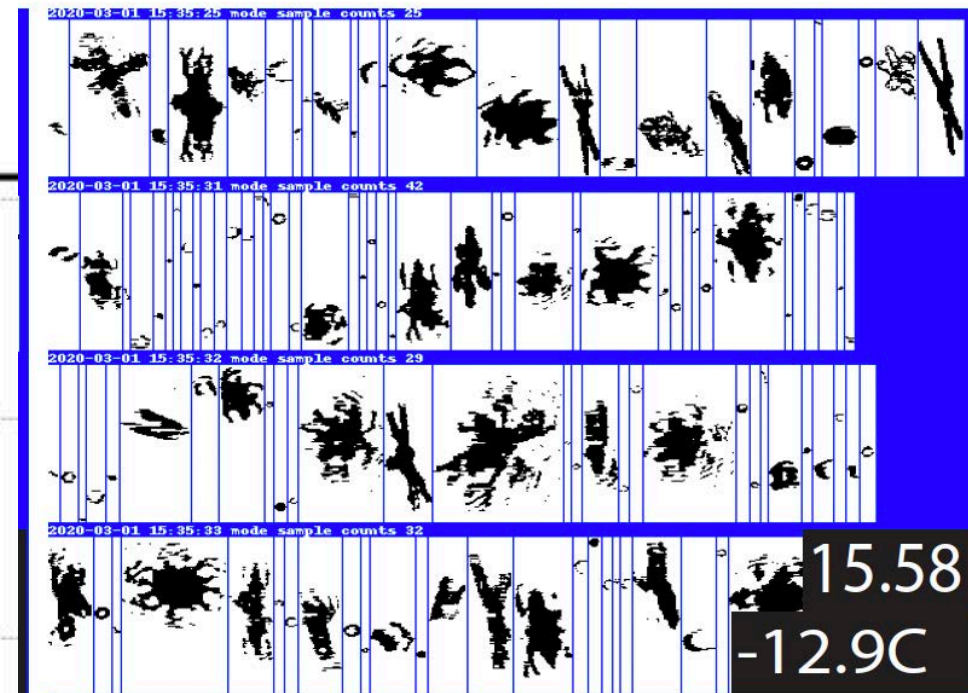
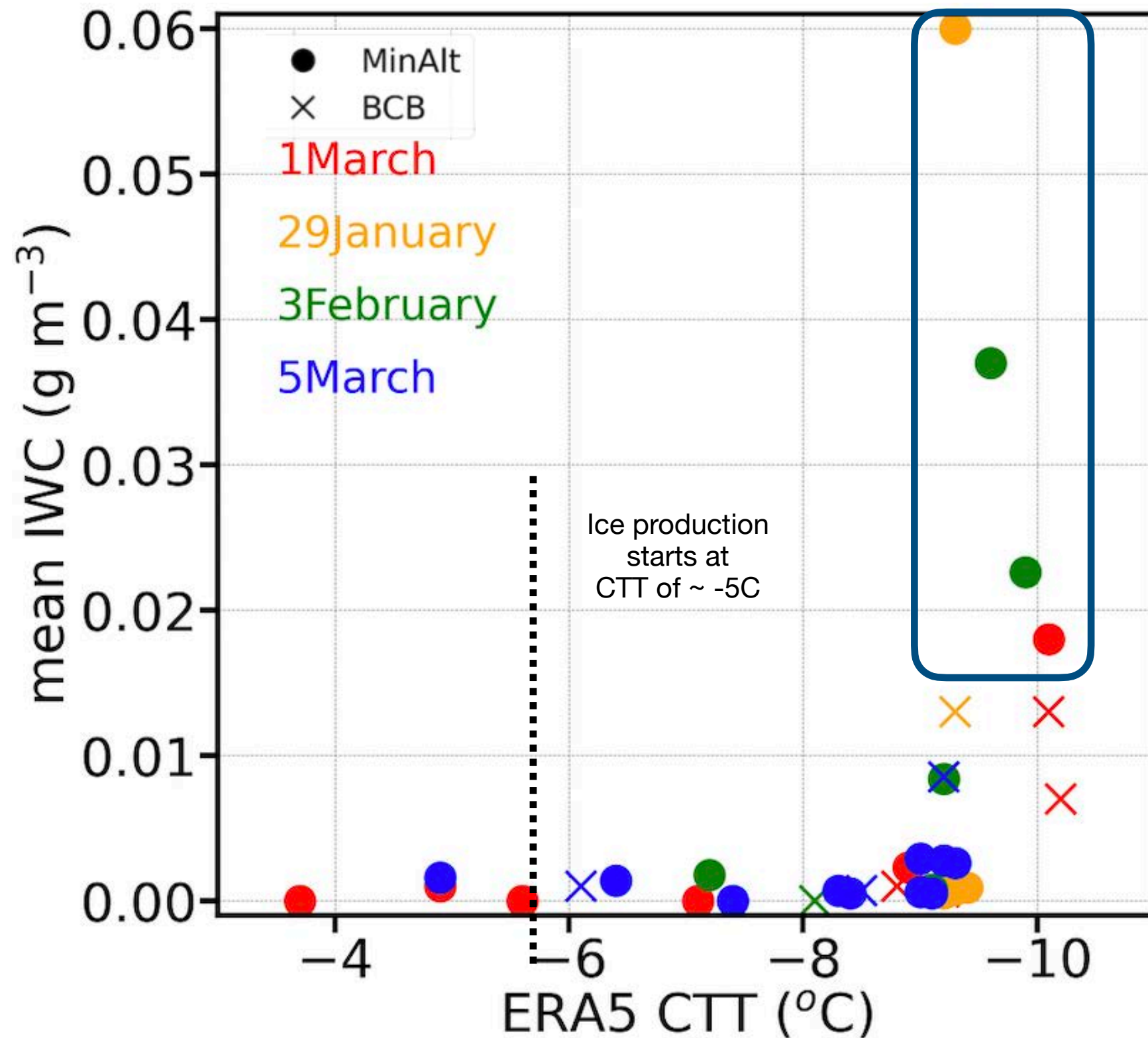
(cloud clearing most likely through  
cloud top entrainment)

# ACTIVATE ice crystal number concentrations indicate secondary ice production, on par with southern ocean values



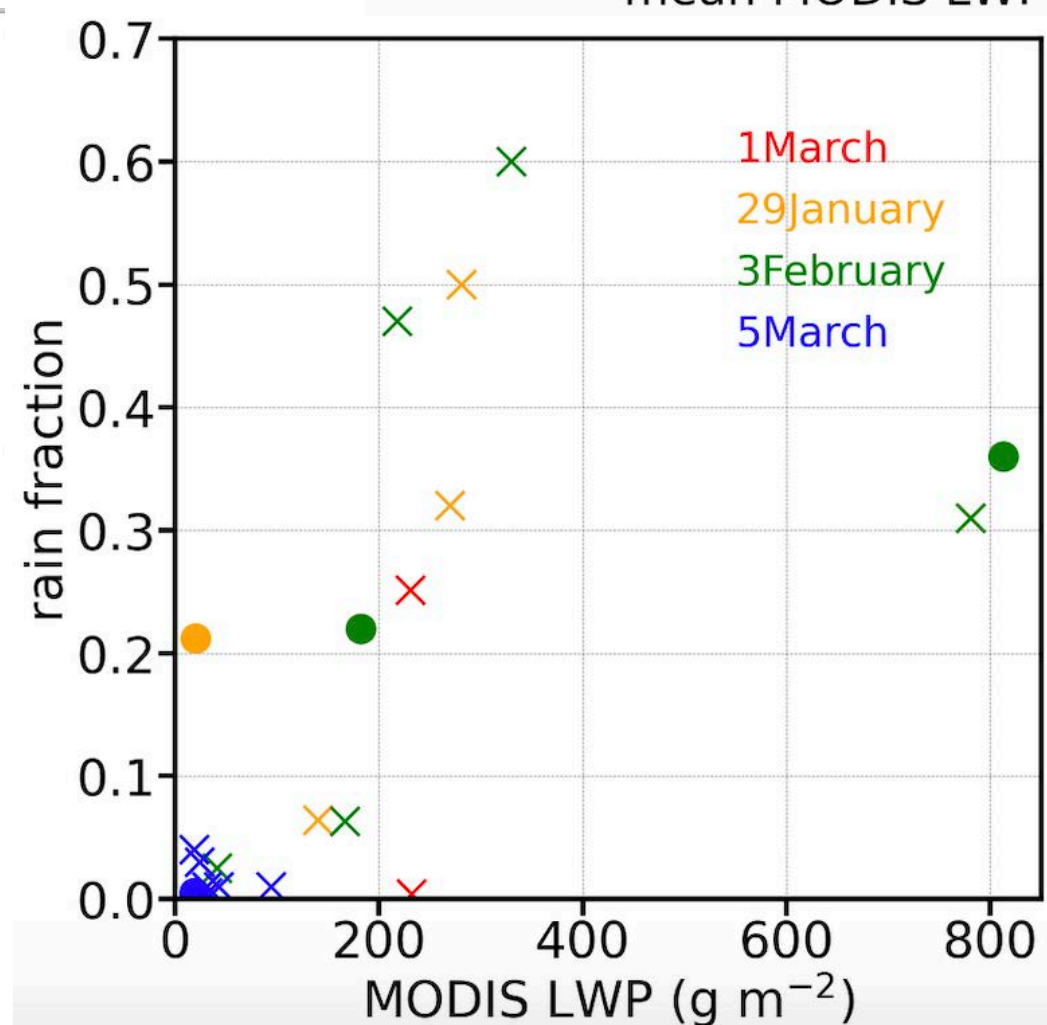
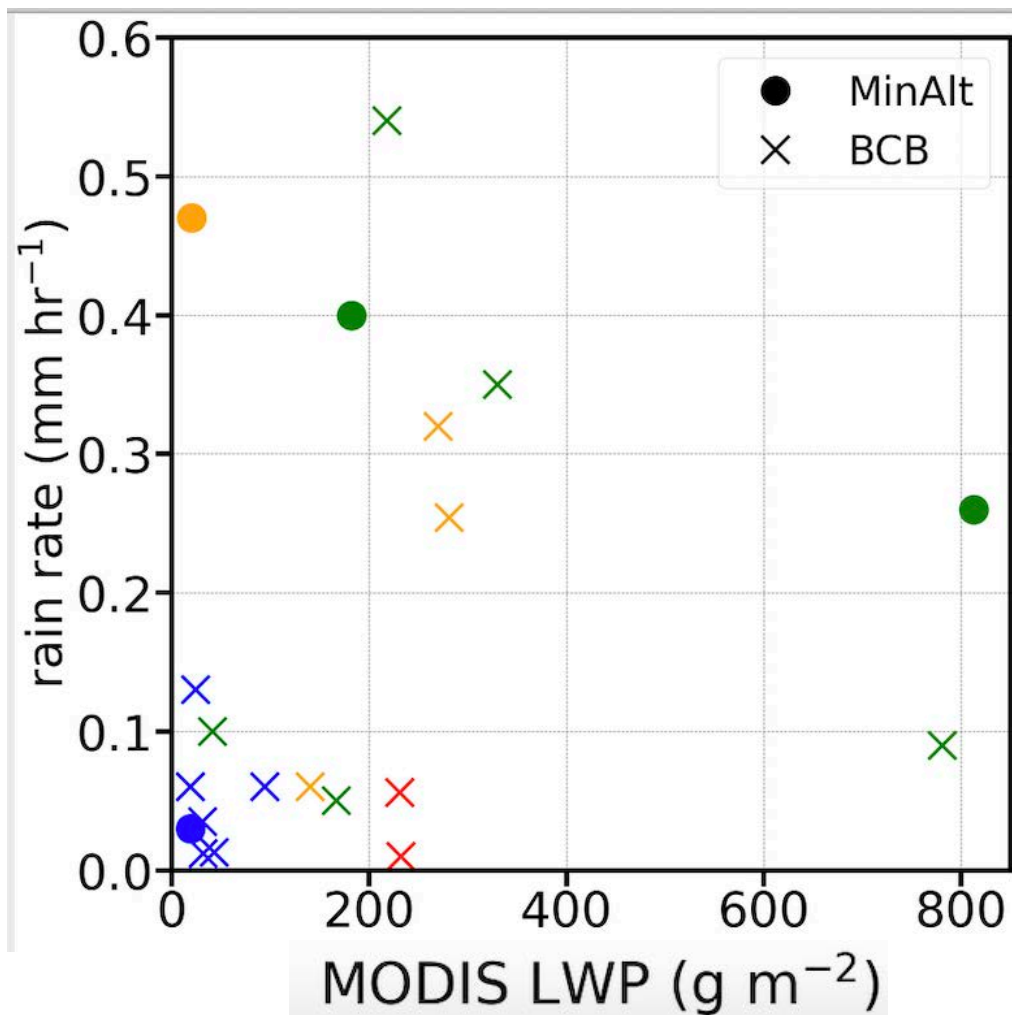
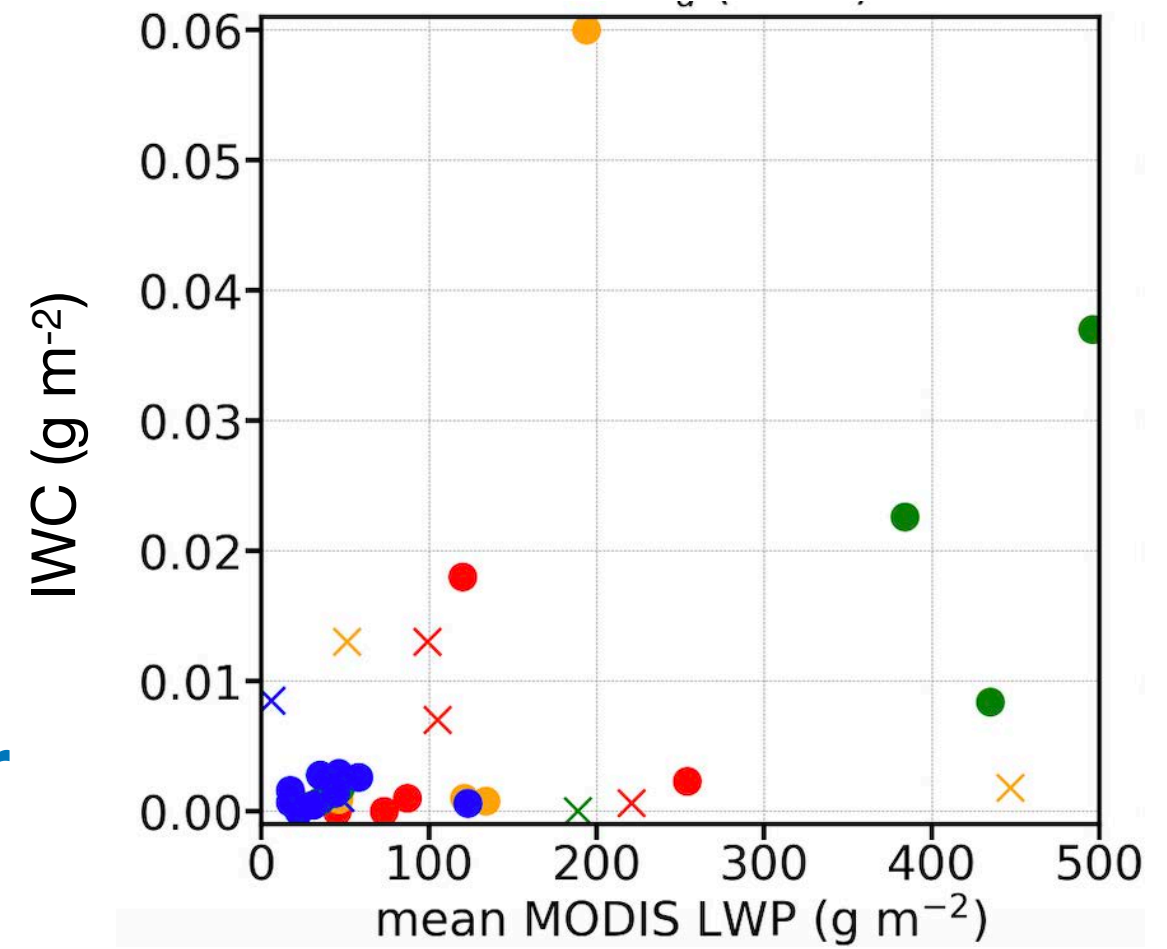


# Ice water contents are higher when vapor diffusional growth dominates, particularly dendritic growth



**Ice water contents are higher for deeper clouds (colder tops), producing more liquid**

**rain (if present) rates are also higher for higher LWPs**





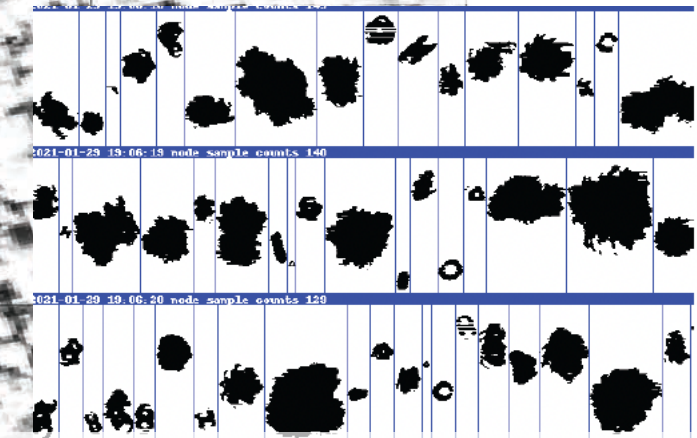
29 January 2021

36°N

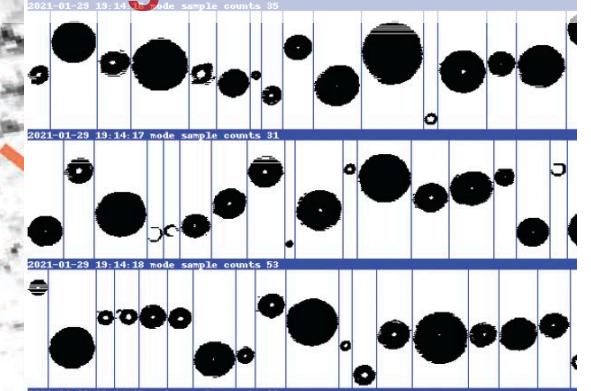
Initially mostly liquid,  
some rimed ice

cloud deepening -> clearer, quicker ice growth  
by vapor diffusion

292K Aggregation



melting into surface rain



294K

250

preceding open-celled  
cloud morphologies

34°N

15.58  
-12.9C

76°W

74°W

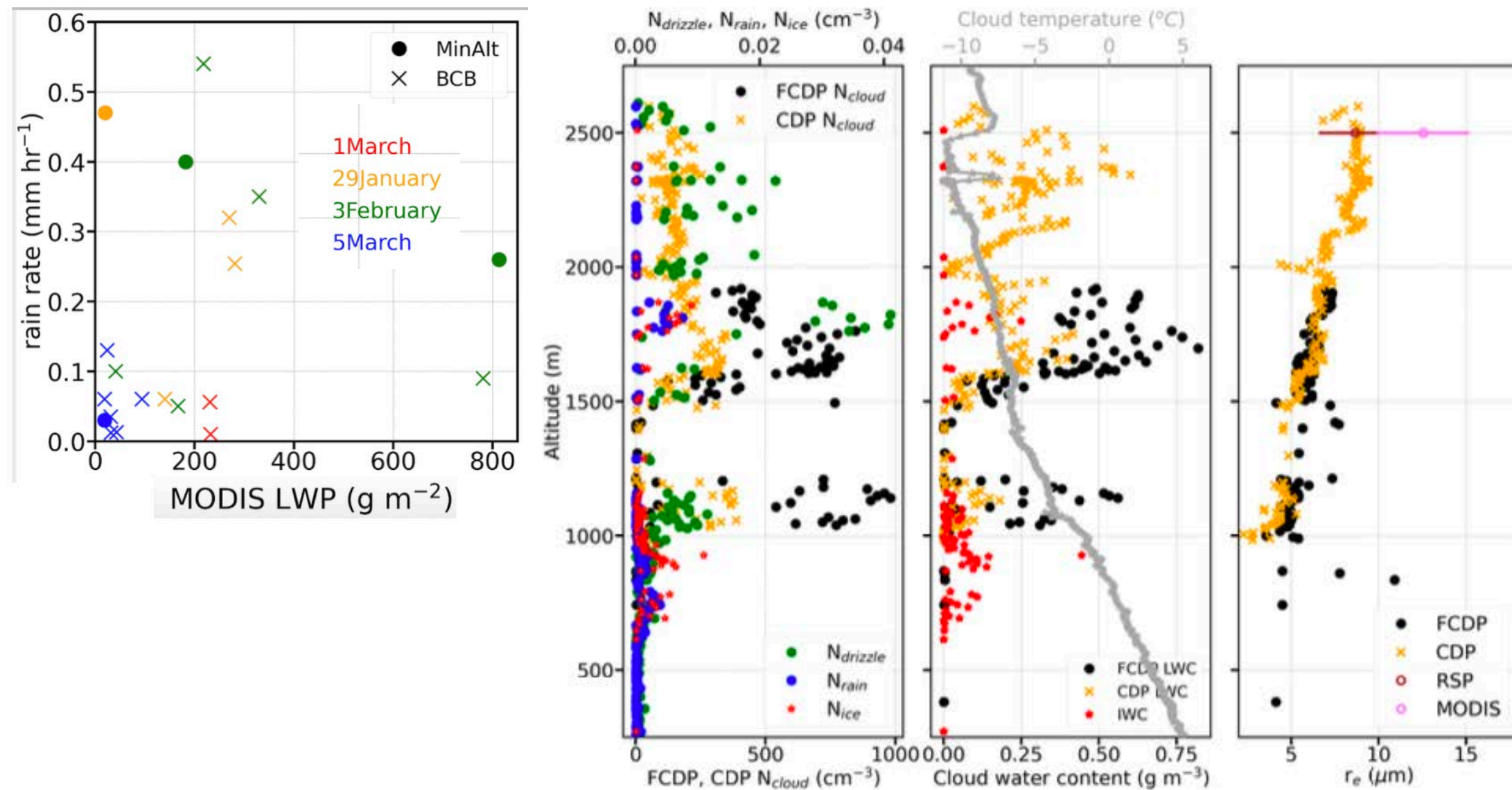
72°W

Cloud depth becomes the primary control on this evolution  
High  $N_d$  almost a given (extending the closed cells)

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Because the MODIS LWPs can reach up to 800 g m<sup>-2</sup> we are also investigating the profiles to put MODIS on a better footing



**Figure 11.** *In situ* February 3, 2021 morning RF44 descent at 33.91°N, 73.07°W profiles. Same conventions as in Fig. 5. FCDP failed in upper cloud layer. CDP  $N_d$  not yet corrected.

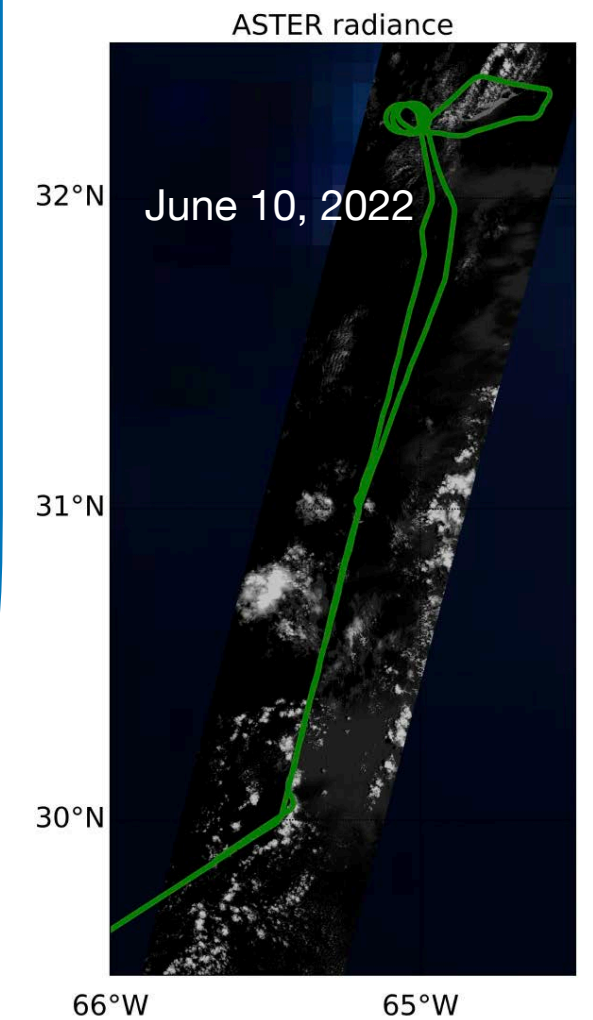
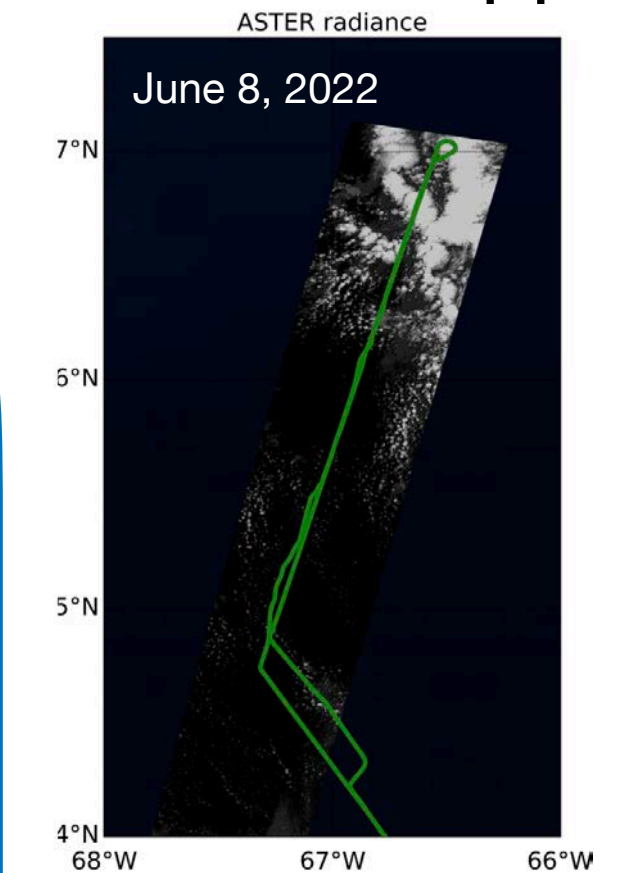
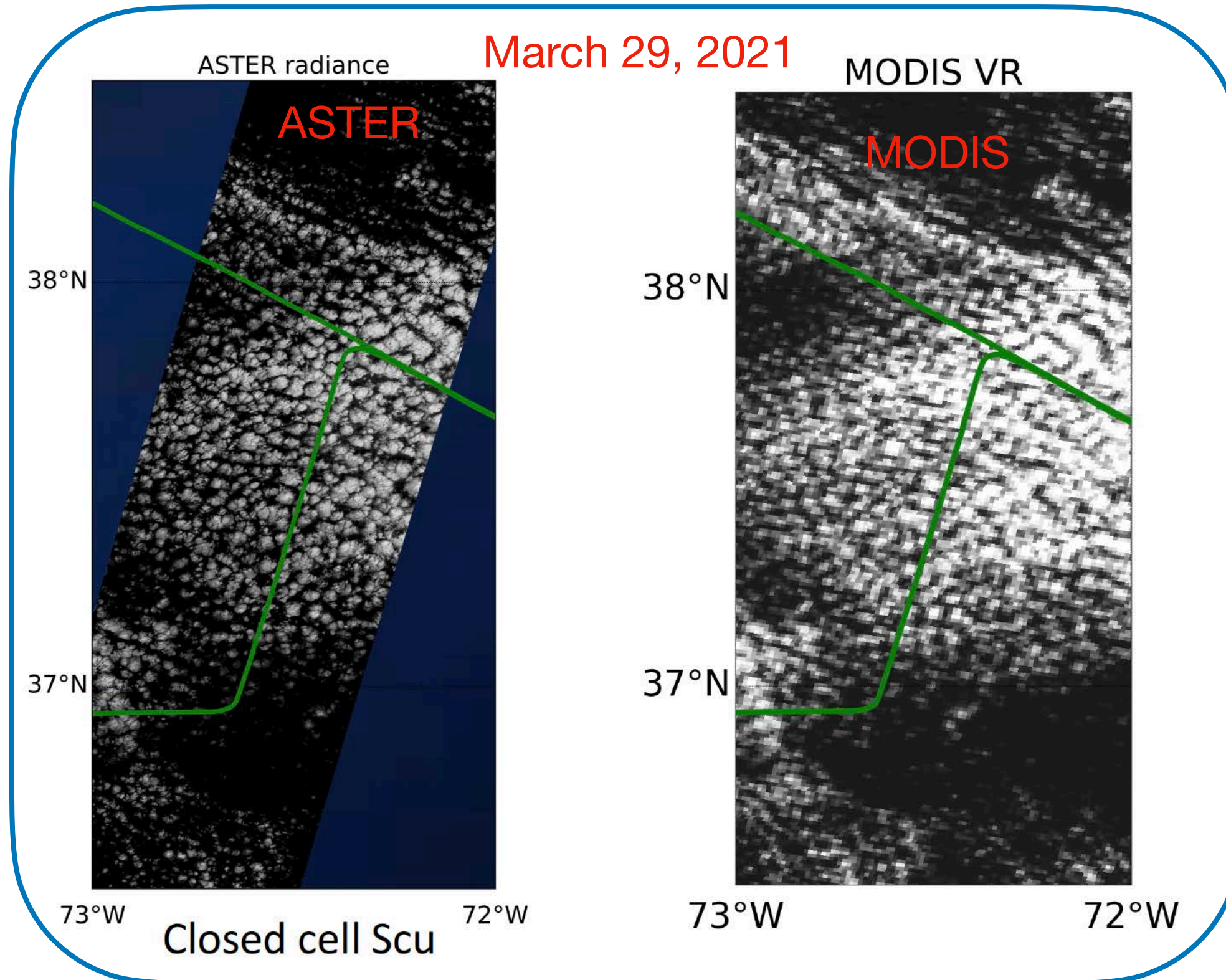


The 5-case CAO analysis is fairly mature.

Plan is to extend the profile analysis to all the wintertime  
ACTIVATE cases as a remote sensing assessment

Shares some goals with the ASTER assessment  
(For which we hope to say something about remotely-sensed  
microphysics as a function of cloud macrophysics)

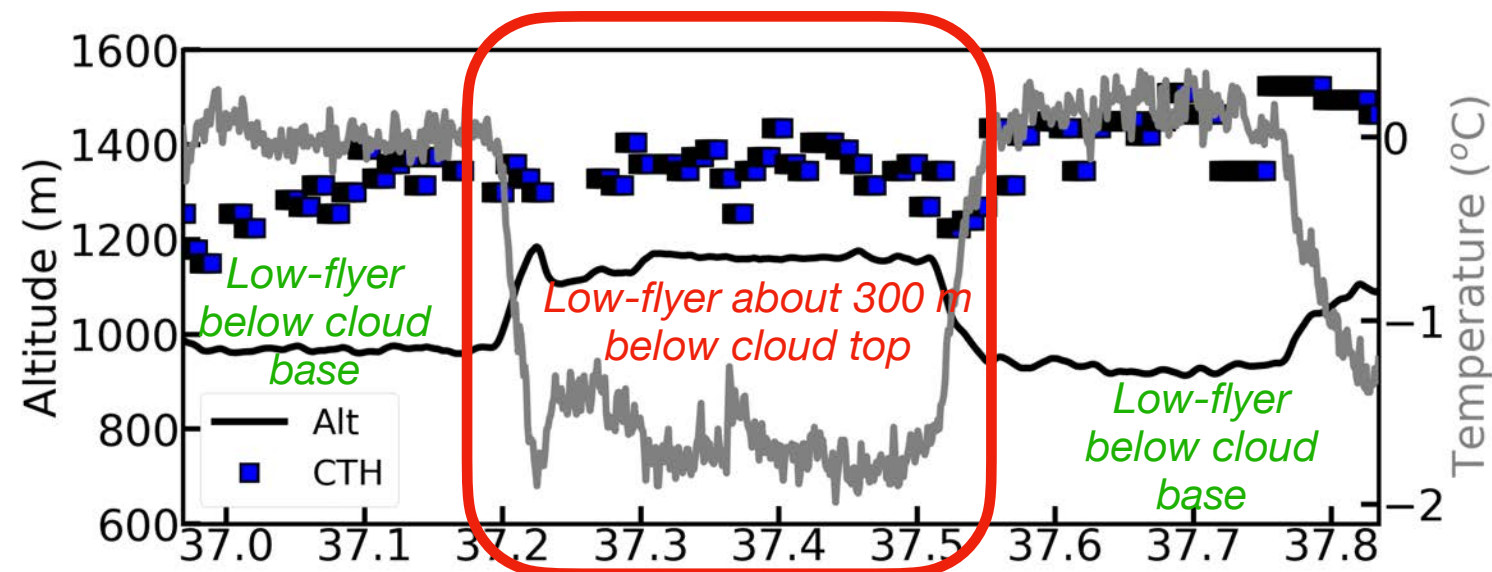
# Three ASTER\* under flights with broken, low clouds & minimum upper-level cloud



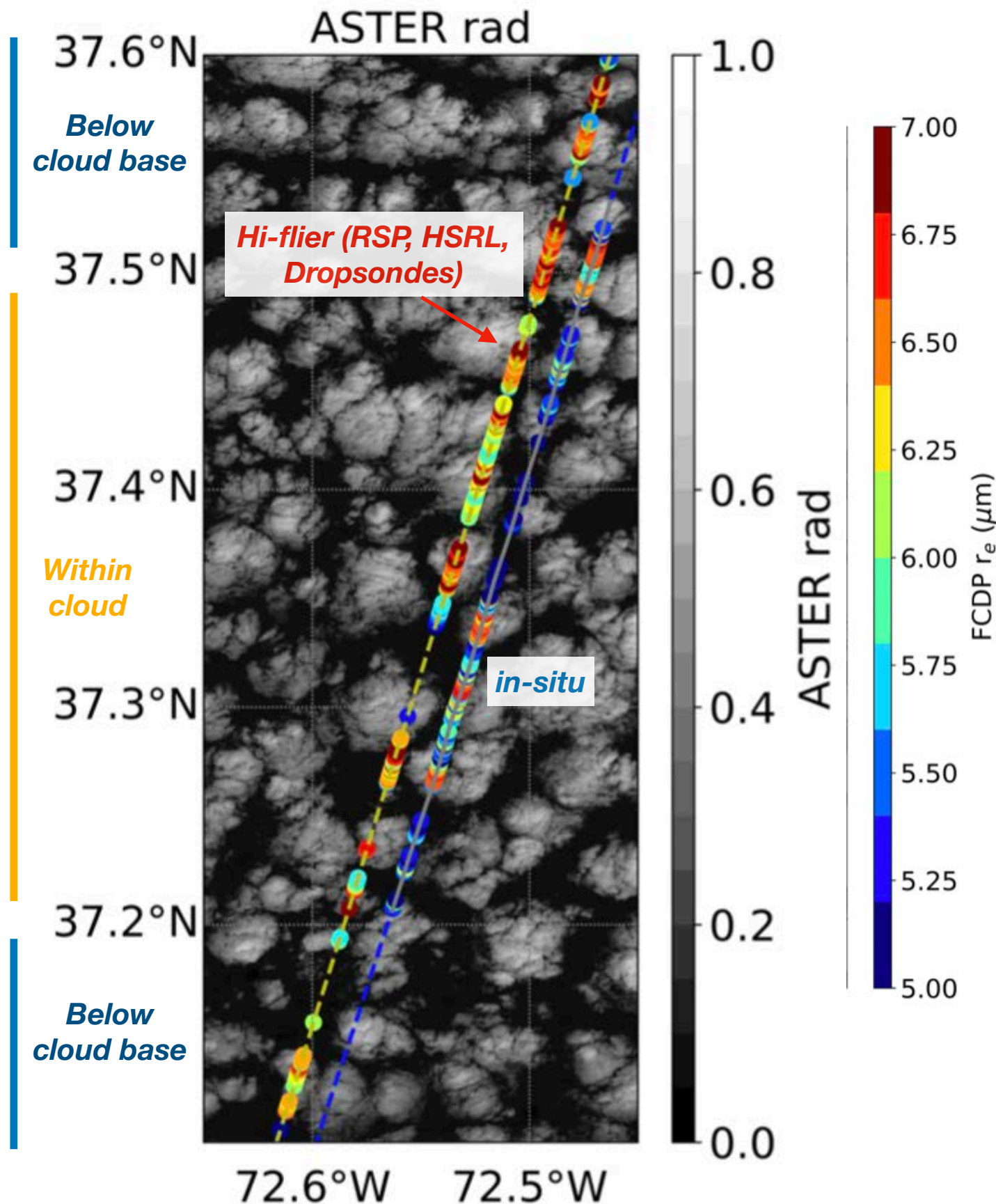
\* Advanced Spaceborne Thermal Emission and Reflection Radiometer



- 15-m spatial resolution cloud mask over a 60 km swath, based on 3 visible nadir bands (0.52- 0.86)+ 11micron (90m resolution)
- No ASTER microphysical retrievals
- RSP derives cloud optical depth ( $\tau$ ), effective radius ( $r_e$ ) using a polarized cloud bow retrieval @ 0.863 micron
- $N_d, LWP$  derived using  $N_d = 1.4067 \times 10^{-6} [cm^{-1/2}] \frac{\tau^{1/2}}{r_e^{5/2}}$  and  $LWP = \rho_w \frac{5}{9} r_e \tau$   
(Painemal and Zuidema, 2011), 'confident cloudy' MODIS cloud mask
- *in-situ*  $N_d$  from FCDP (1.5-50 micron) and CDP



March 29, 2021, 15:37-15:50 UTC  
~50 km span

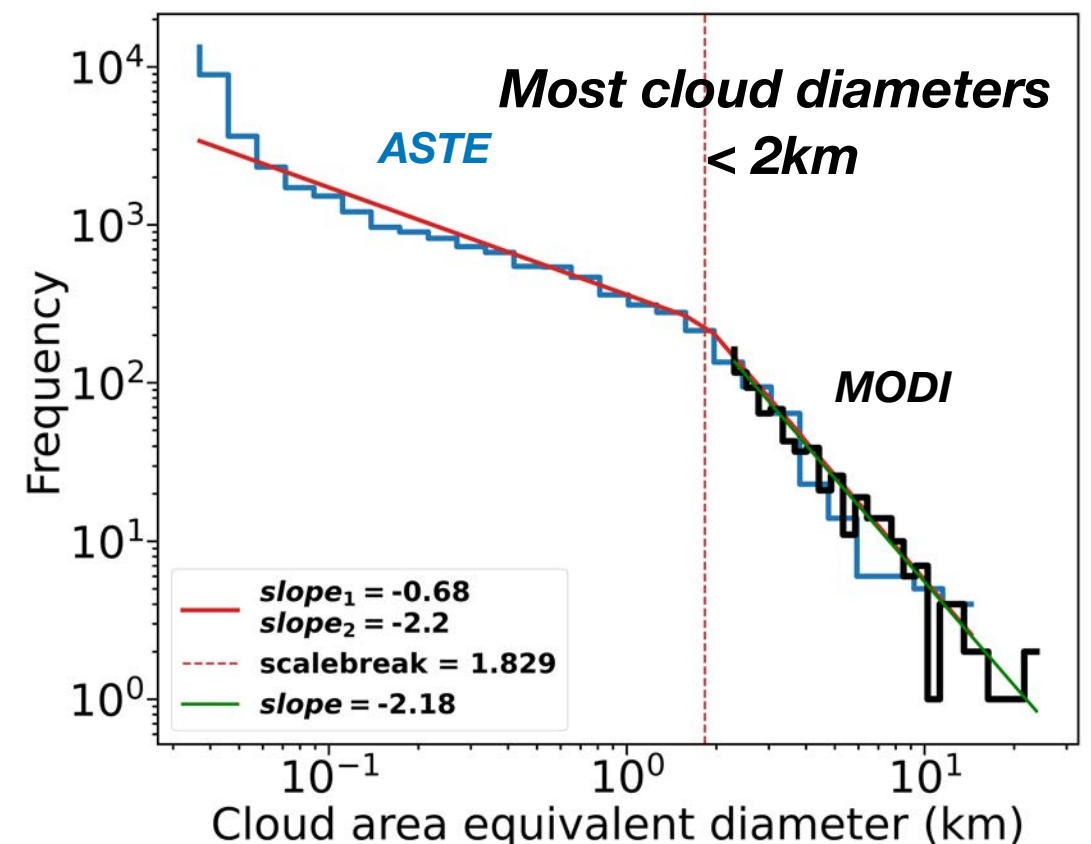


Includes in-situ corrected time offset; 5 minutes diff from ASTER,  
2km offset between the two planes

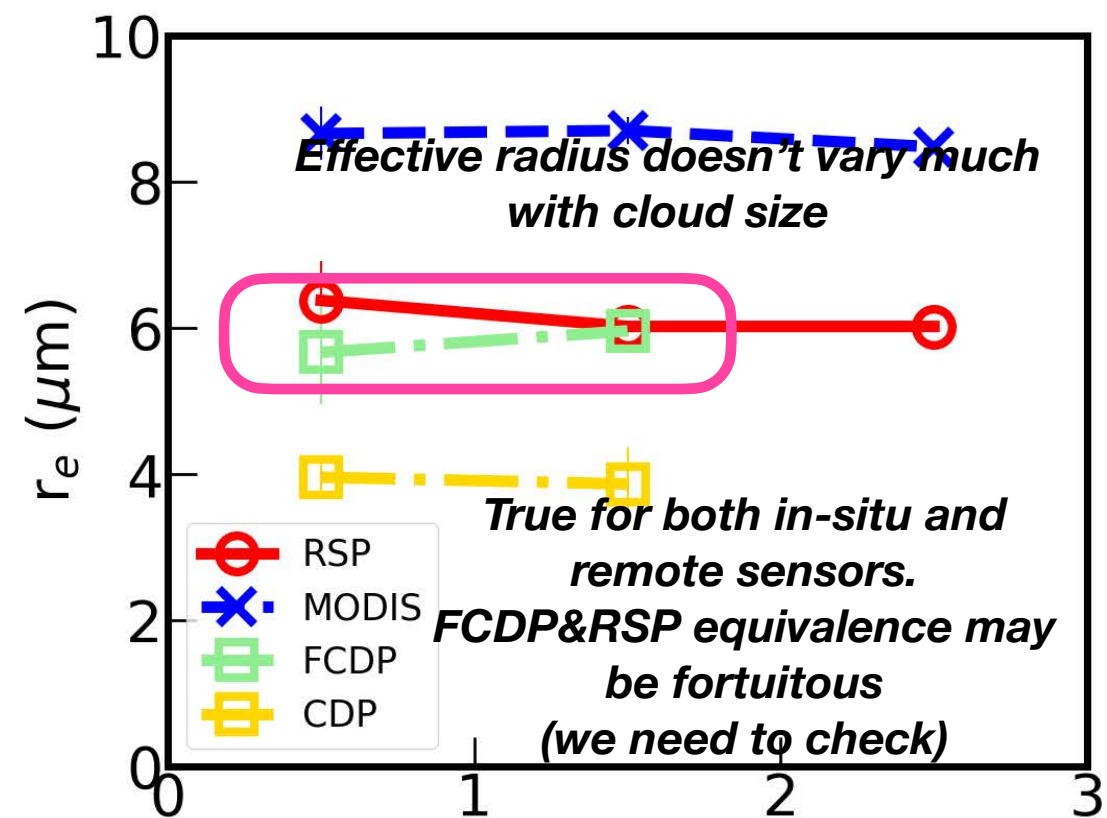
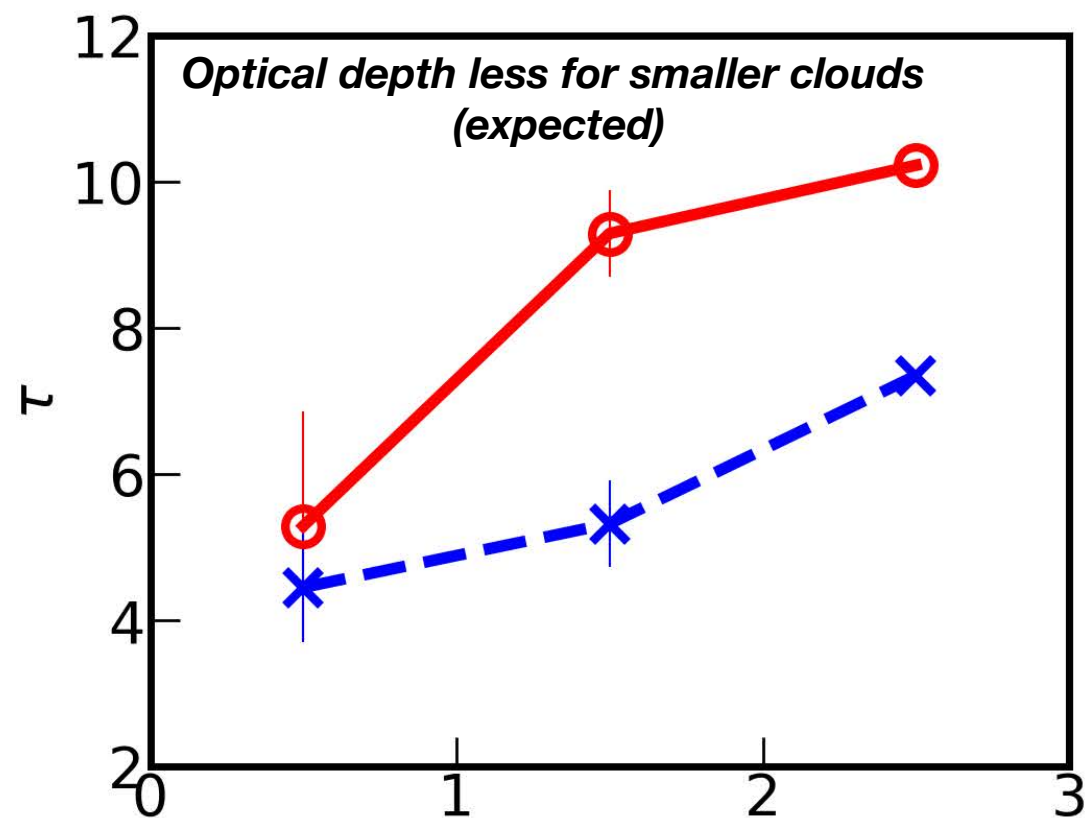
Not a perfect colocation but: -  
a homogeneous aerosol &  
cloud environment:

- $N_{aerosol}$  ( $>100$  nm) of  
200-250  $\text{cm}^{-3}$
- $H_{ct}$  of 1.2-1.4 km,  $T_{ct} > -10$   
 $^{\circ}\text{C}$
- Little (no?) rain

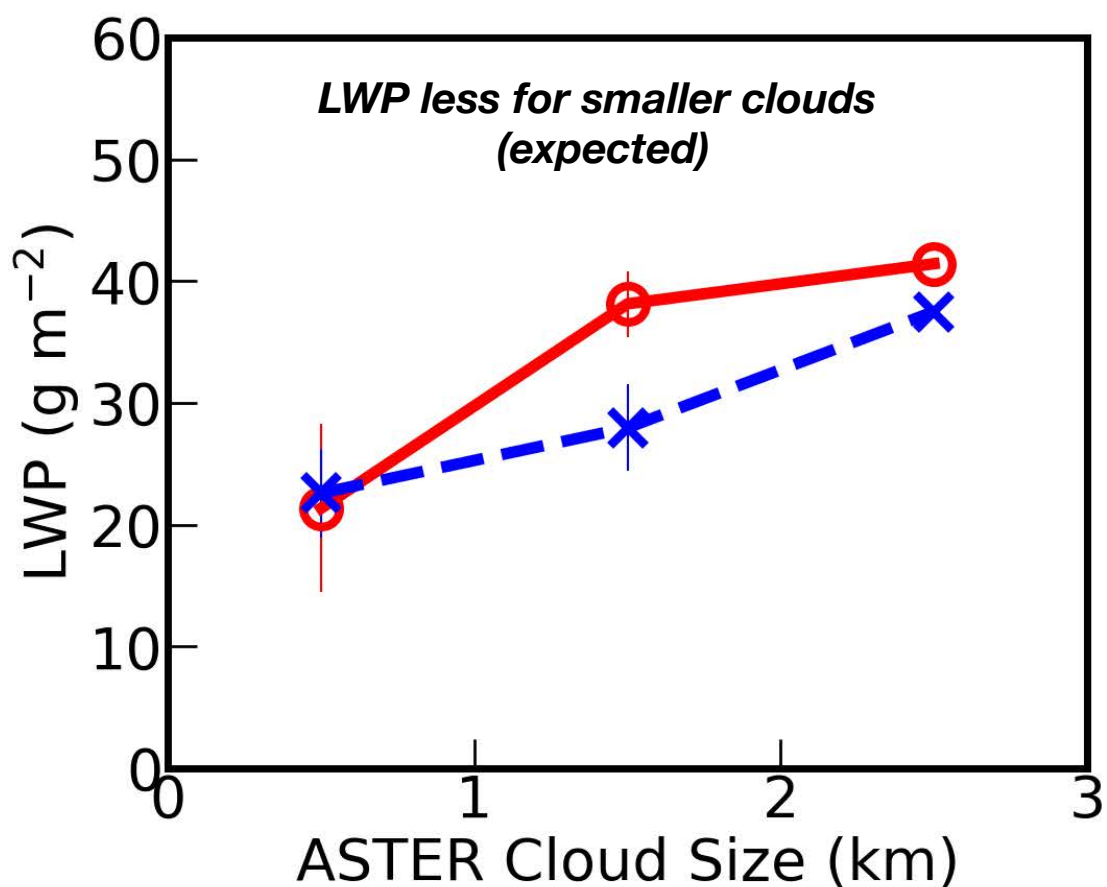
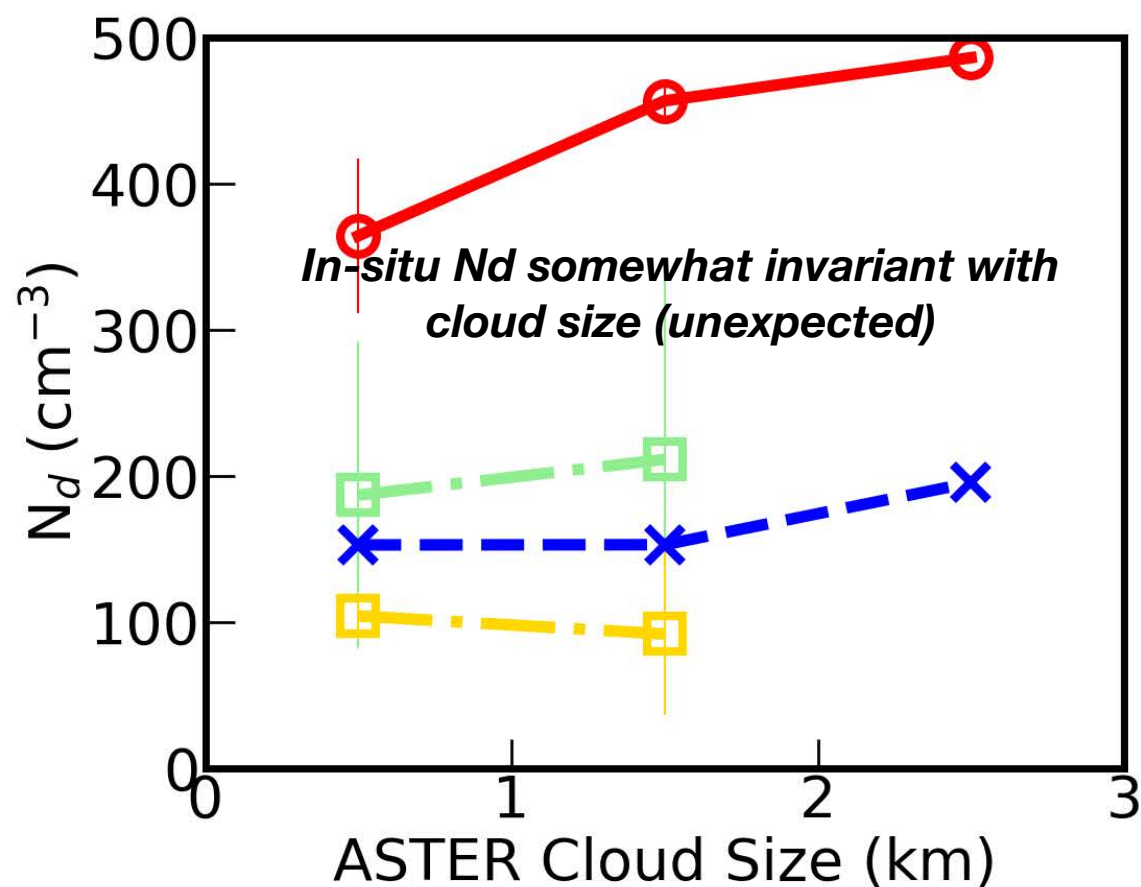
=> statistical relationships  
okay







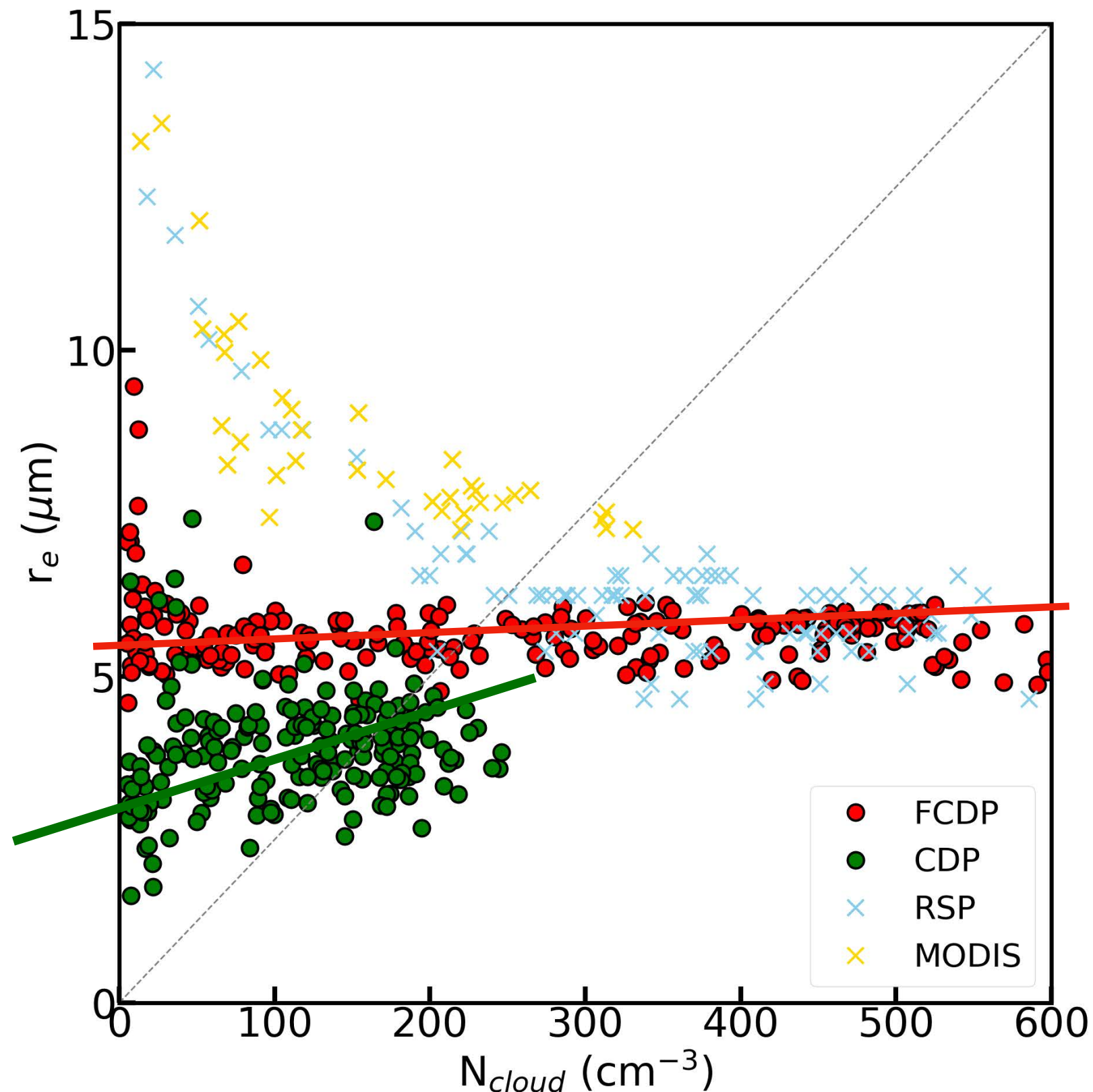
*Based on 5 10-km domains*



In-situ probes show  $N_d$ ,  $r_e$  are [slightly] positively correlated, suggesting growing inner cloud cores & dissipating edges [but still need to put that on firmer ground]

While remote retrievals suffer at the edges, RSP extends to smaller cloud sizes/higher  $N_d$

Good correspondence between RSP, FCDP  
Ignores 300m of vertical Displacement





can UC-12 camera imagery be used to extend to more scenes (10 m spatial resolution, ~15 km swath)?

Unclear, might at the least need to identify some useful cases

